

## Identification of possible tsunami earthquakes along the Mexican Subduction Zone

○Ketzallina FLORES Yoshihiro ITO, Emmanuel Soliman M. GARCIA

Tsunami earthquakes are defined as events that generate tsunamis of much greater amplitude than those expected from their seismic magnitudes (Kanamori 1972), and the source process is still controversial. This is because tsunami earthquakes are relatively rare and have not been observed yet with recent seismic and geodetic networks.

On April 18<sup>th</sup>, 2002, the earthquake with  $M_w$  6.7 was located about 55 km from the coast of Guerrero, Mexico. The earthquake occurred near the trench of the northwest Guerrero seismic gap, which has not broken in a large earthquake since 1911 (Singh et al., 1981; UNAM Seismology Group 2015). The rupture of the 2002 event most likely began at a subducted seamount, and propagated unilaterally towards NW, parallel to the trench for  $\sim 60$  km and for a duration of  $\sim 60$  s. The moment rate function was highly rugged with two dominant pulses separated by about 50 s. Although relatively small in magnitude, the earthquake has all the characteristics of a tsunami earthquake: the slip occurs very close to the trench; the rupture speed is slow ( $\sim 1$  km/s); the high-frequency radiation is deficient, the moment-scaled radiated energy is lower ( $E_s/M_0 = 1.43 \times 10^{-6}$ ) than those of typical interplate earthquakes. This event was also characterized by a large centroid delay time from the origin time, which was about 30 s (Duputel et al., 2013). Also, it produced anomalously low accelerations and generated a small tsunami (Iglesias et al., 2003). Therefore, this event can be categorized as a tsunami earthquake.

Okal and Borrero (2011) conducted a detailed seismological study of the large Colima, Mexico earthquake. The event occurred on June 3<sup>rd</sup> and its

aftershocks on June 18<sup>th</sup> and 22<sup>nd</sup> in 1932. The aftershock on June 22<sup>nd</sup> generated a tsunami more devastating than that of the main shock, despite much smaller seismic magnitudes. They confirmed a deficient energy-to-moment ratio derived from high-frequency P-waves recorded in Pasadena, approximately  $E_s/M_0 = 6.607 \times 10^{-7}$ . Their study suggests that this aftershock had the characteristics of a tsunami earthquake.

Therefore, in the previous two examples, we have observed that the moment-scaled radiated energy ( $E_s/M_0$ ) plays a fundamental role in characterizing a tsunami earthquake. Newman and Okal (1998) also demonstrated that the scaled energy is a powerful discriminant for tsunami earthquakes.

Shapiro et al. (1998) apply a method based on the ratio of the total radiated energy to the high-frequency energy, ER, computed from at the broadband seismometer at CUIG, located in Ciudad Universitaria, Mexico City. Using this information, Iglesias et. al., (2003) showed a relation between ER and  $E_s/M_0$ , which suggests that events with high values of ER should also have low values of  $E_s/M_0$ .

In general, if ER is greater than 100 for an event, then  $E_s/M_0$  is less than  $3 \times 10^{-6}$ . For instance,  $E_s/M_0$  values were calculated as  $1.5 \times 10^{-6}$ ,  $0.6 \times 10^{-6}$ , and  $2.6 \times 10^{-6}$  for the tsunami earthquakes of Nicaragua (2 September 1992,  $M_w$  7.6), Java (2 June 1994,  $M_w$  7.8), and Peru (21 February 1996,  $M_w$  7.5), respectively (Venkataraman, 2002). ERs for two tsunami earthquake were 1180 and 512 in Peru in 2002 (Iglesias et al., 2003).

The seismic magnitude scales of moment magnitude,

$M_W$  and surface wave magnitude,  $M_S$ , body wave magnitude,  $m_b$  have a fundamental role in identifying a size of earthquakes and expecting a scale of tsunamis. A tsunami earthquake is generally identified as an event with a larger  $M_W$  than ordinary earthquakes with the same  $M_S$ , whereas  $M_S$  usually saturates when a very large earthquake has an ordinary source process (Polet and Kanamori, 2000).

In this work, we identify possible tsunami earthquakes that occurred in the Mexican subduction zone. First, we select the events with ER greater than 100. Next, we compare  $E_S/M_0$  of some possible tsunami earthquakes with ordinary and tsunami earthquakes. Finally, we use the comparisons of  $M_S$ ,  $m_b$ ,  $M_W$  to determine if these events share common characteristics with tsunami earthquakes. It was previously suggested in a study using bathymetric data (Geersen et al, 2019) that lower-plate topography and sediment thickness may be related to the risk of a tsunami earthquake. Therefore, we will also compare the location of the possible tsunami earthquake events with residual gravity values in the surrounding area to help understand any possible correlation between the source of the events and the geological structure.

## References

Duputel, Z., Tsai, C.V., Rivera, L. & Kanamori, H., 2013. Using centroid time-delays to characterize source durations and identify earthquakes with unique characteristics. *Earth Planet. Sci. Letter.* <https://doi.org/10.1016/j.epsl.2013.05.024>.

Geersen, J., 2019. Sediment-starved trenches and rough subducting plates are conducive to tsunami earthquakes. *Tectonophysics*, 762, pp. 28-44.

Iglesias A., Singh S.K., Pacheco J.F., Alcántara L., Ortiz M. & Ordaz M., 2003. Near-trench Mexican earthquakes have anomalously low peak accelerations, *Bull. Seism. Soc. Am.*, 93, 2, 953–

959.

Kanamori, H., 1972. Mechanism of tsunami earthquakes. *Physics of the Earth and Planetary Interiors*, Volume 6, Issue 5, 346-359, ISSN 0031-9201, [https://doi.org/10.1016/0031-9201\(72\)90058-1](https://doi.org/10.1016/0031-9201(72)90058-1).

Newman, A. V., and E. A. Okal, 1998. Teleseismic estimate of radiated seismic energy: the  $E_S/M_0$  discriminant for tsunami earthquakes, *J. Geophys. Res.* 103, 26,885–26,898.

Okal, E.A., Borrero, J.C., 2011. The ‘tsunami earthquake’ of 1932 June 22 in Manzanillo, Mexico: seismological study and tsunami simulations, *Geophysical Journal International*, Volume 187, Issue 3, Pages 1443–1459.

Polet, J., and Kanamori, H., 2000. Shallow subduction zone earthquakes and their tsunamigenic potential. *Geophysical Journal International* 142, 684-702.

Shapiro, N. M., S. K. Singh, and J. F. Pacheco, 1998. A fast and simple diagnostic method for identifying tsunamigenic earthquakes, *Geophys. Res. Lett.* 25, 3911–3914.

Singh S.K., Astiz, L. & Havskov, J., 1981. Seismic gaps and recurrence periods of large earthquakes along the Mexican subduction zone: A reexamination, *Bull. Seismol. Soc. Am.*, 71, 827-843.

Singh, S.K., Arroyo, D., Pérez-Campos, X., Rodríguez, Q., Iglesias, A. & Ortiz, M., 2016. Fast identification of near-trench earthquakes along the Mexican subduction zone based on characteristics of ground motion in Mexico City, *Bull. Seism. Soc. Am.*, 106, 2071–2080. <https://doi.org/10.1785/0120160003>.

UNAM Seismology Group (with contribution from UAM, Azcapotzalco, Mexico. D.F.), 2015. Papanaoa, Mexico earthquake of 18 April 2014 (MW 7.3), *Geofísica Internacional*, 54(4), 363-386.