

The Second Global Summit of Research Institutes for Disaster Risk Reduction  
Development of a Research Road Map for the Next Decade  
March 19-20, 2015

# New Conceptual Model of Large-Scale Landslide

Reported by

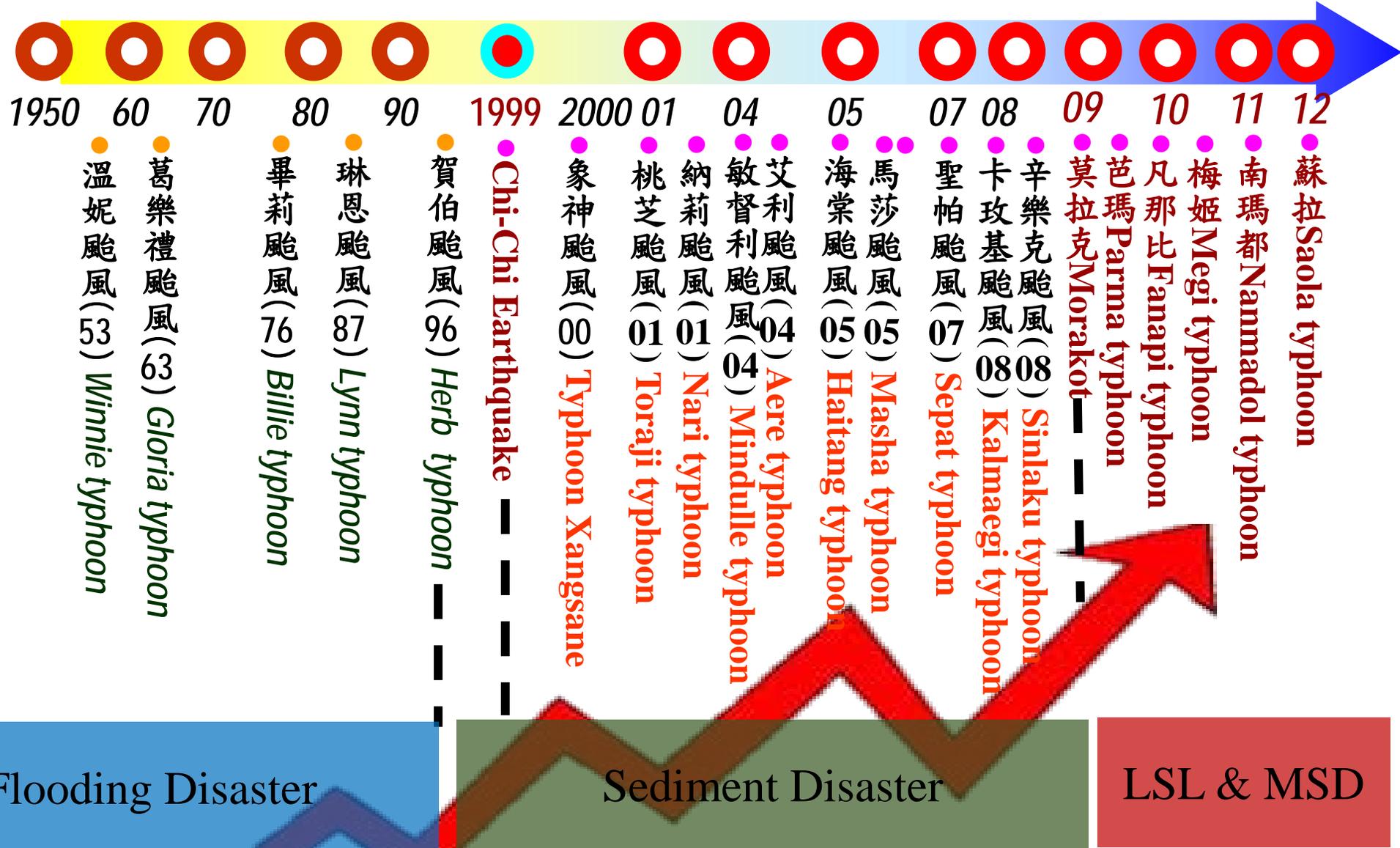
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Director, Disaster Prevention Research Center

National Cheng Kung University



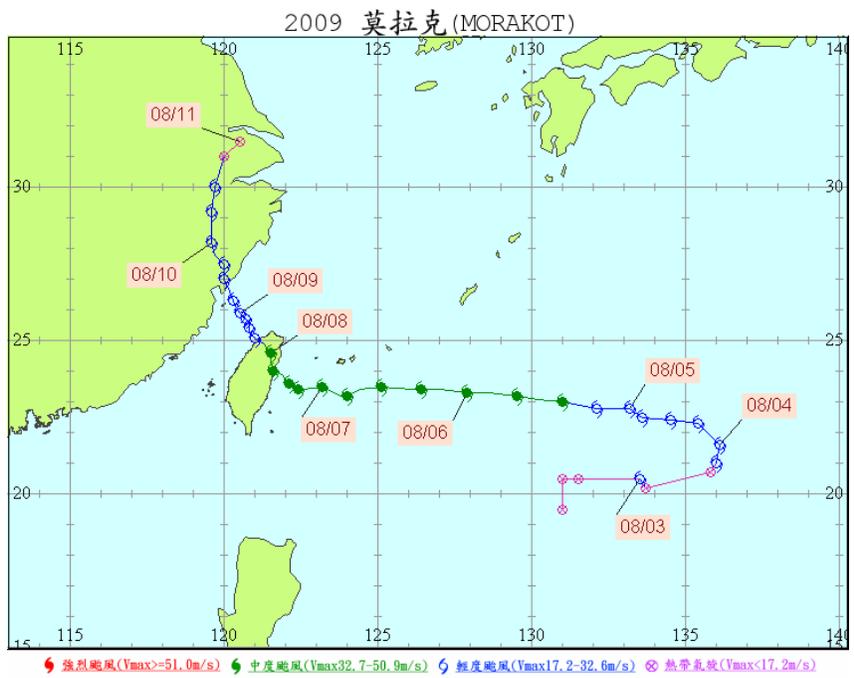
To review historical disasters from 1950, **LSL & MSD** occur frequently.



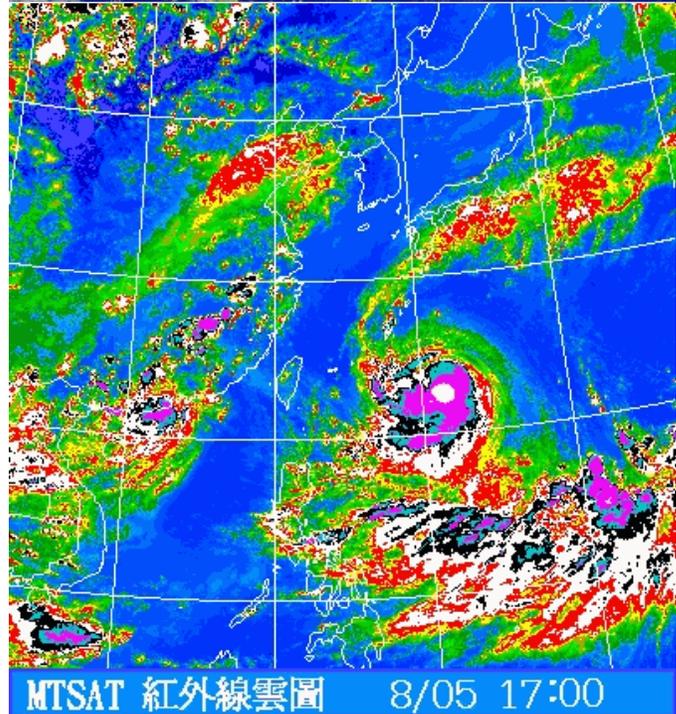
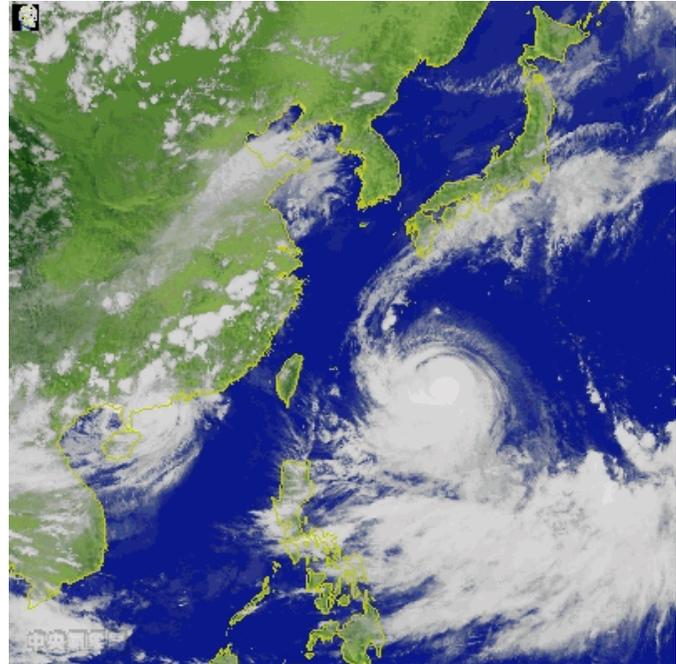
Background
Case Study
Idea
Experiment
Modelized
Solution
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Conclusion

# Typhoon Morakot

- The storm turned into a typhoon on Aug. 4.
- Rainfall started on Aug. 6.
- The eye of the typhoon left Taiwan from Taoyuan at 14:00 on Aug. 8.
- The rainfall continued until Aug. 10.



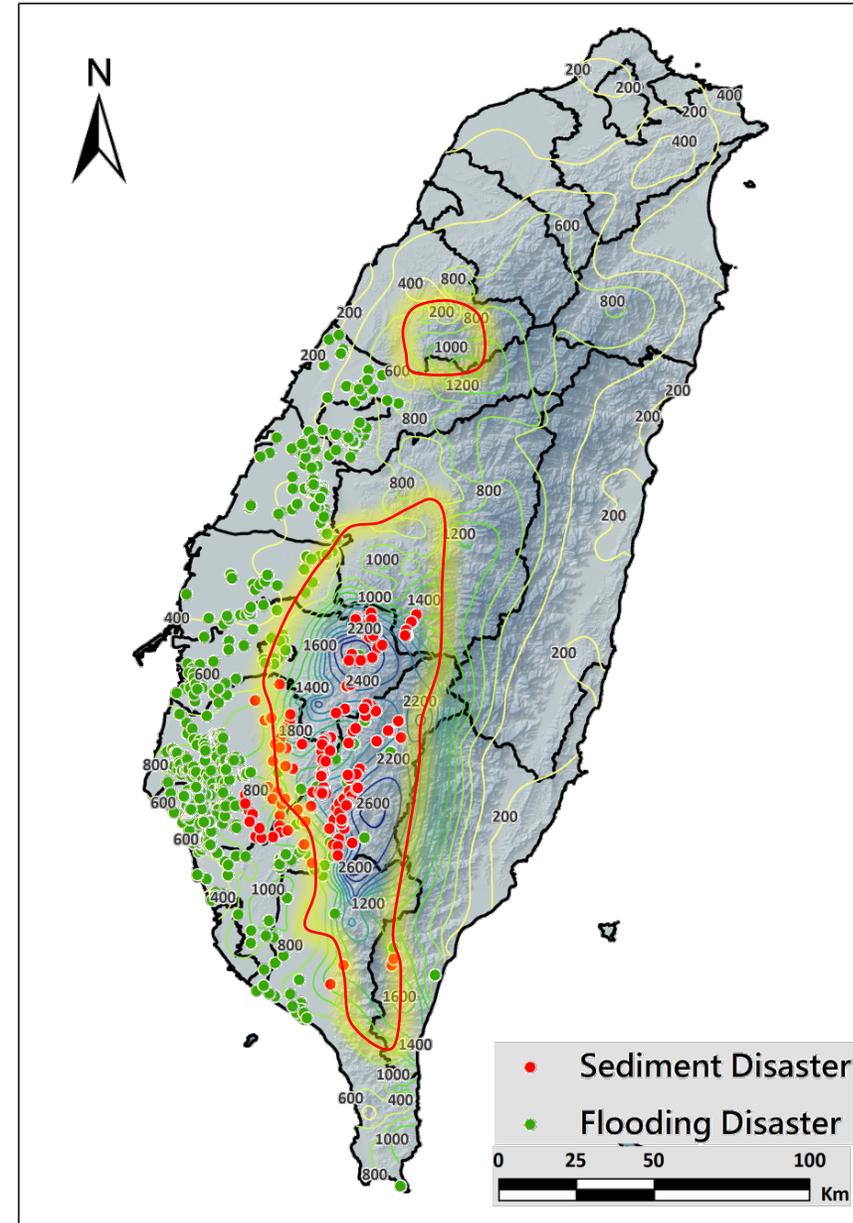
- Path of the center of Typhoon Morakot



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# Disasters of Typhoon Morakot

- **Alisan Station**
  - Long duration (91 hours)
  - High intensity (123 mm/hour)
  - Large accumulated rainfall depth (3000 mm-72 hour)
  - Broad extent (one-fifth of Taiwan was covered)
- **Sediment-related disasters in the mountain area located within the range of precipitation > 1,000mm**
- **Inundation area located at the downstream of rainfall center in the plain area**



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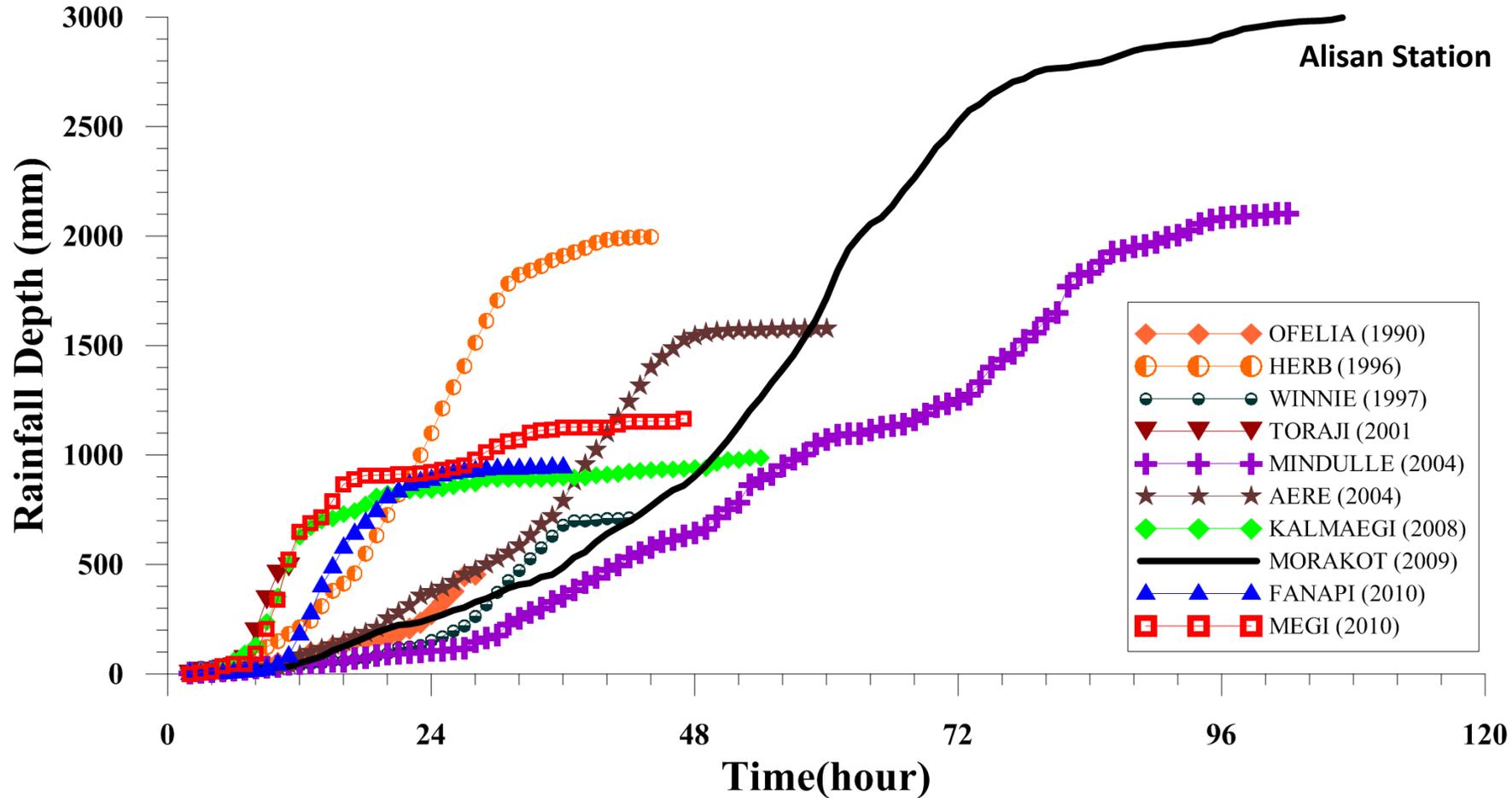
Solution

Comparison

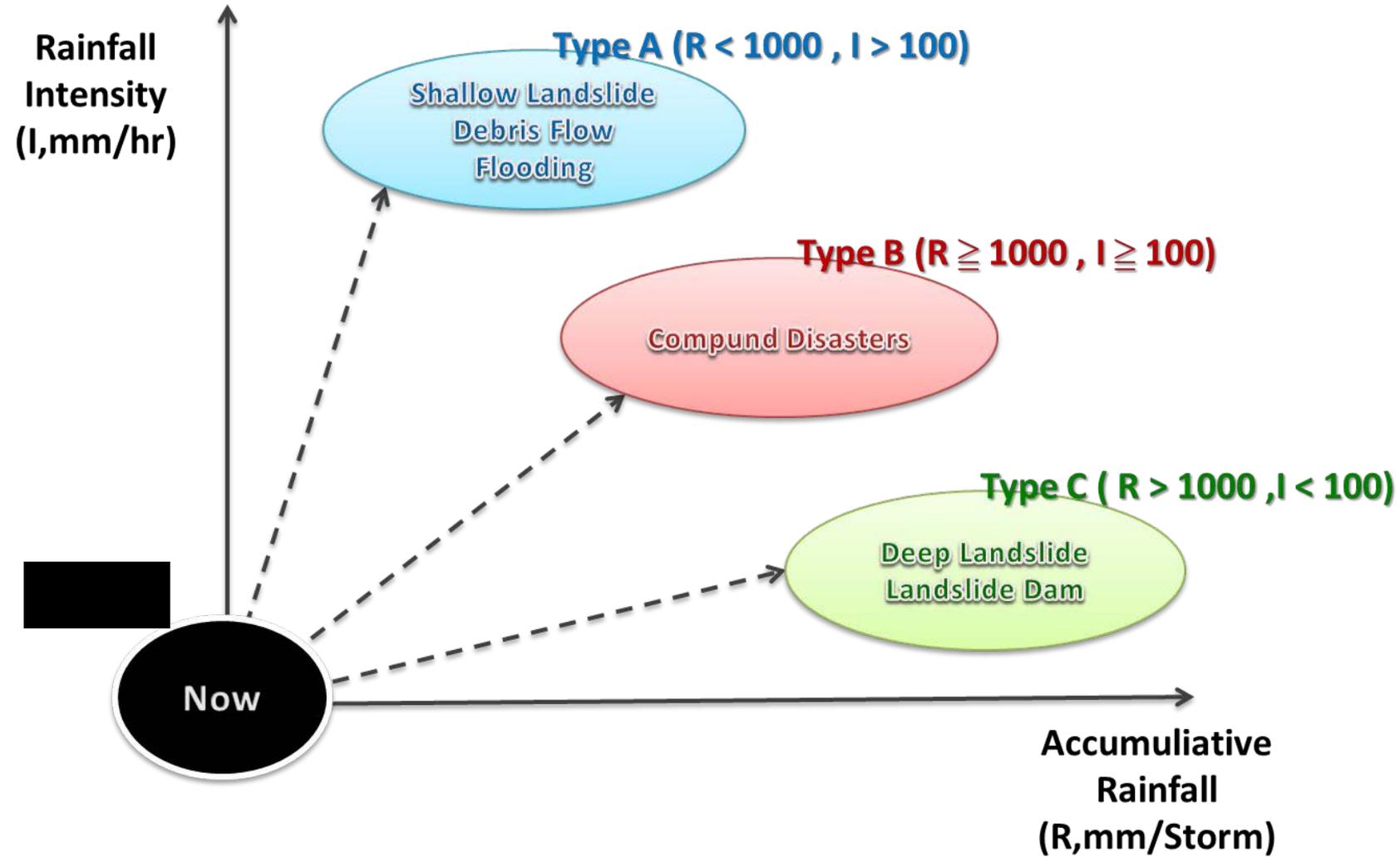
Conclusion



# The relationship between accumulative rainfall and duration of catastrophic typhoons



- Background
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• Shieh, 2012

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# The scale and type of the disaster increasing with the frequent appearance of extreme weather

Landslide & Debris Flow disaster



Debris flow disaster



Natural Dam



Natural Dam



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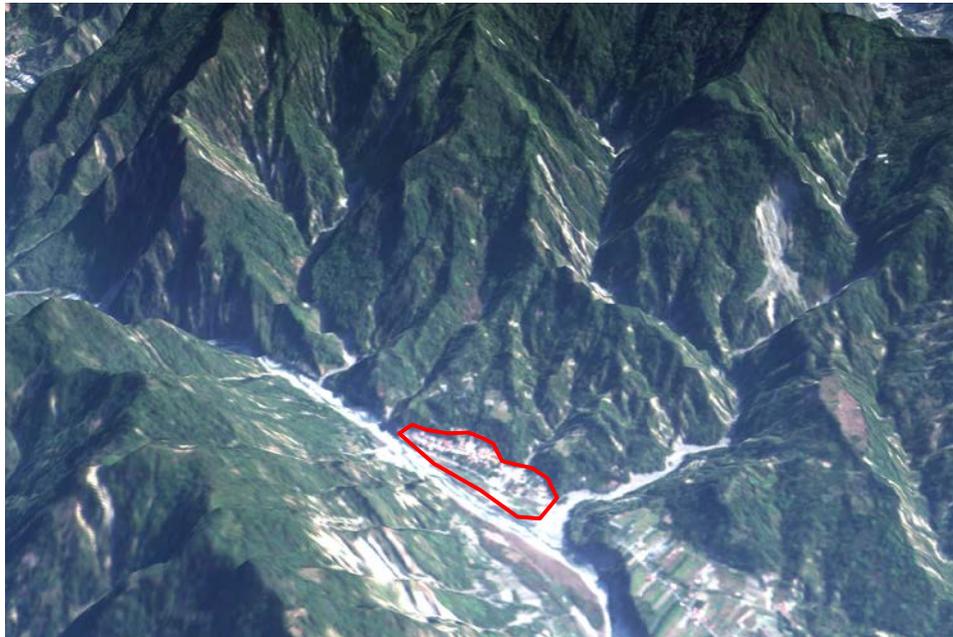
Solution

Comparison

Conclusion

# Large-scale landslide and compound disaster become a new challenge

- Area : 202 ha Depth : 80 meter Volume : 24 million m<sup>3</sup>



**Before Typhoon Morakot in Hsiaolin Village  
(FORMOSAT-2)**



**After Typhoon Morakot in Hsiaolin Village  
(FORMOSAT-2)**

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## LESSONS FROM THE DISASTER OF HSIAOLIN VILLAGE

- **Compound Disaster, included shallow landslide, debris flow, flooding, large-scale landslide, natural dam break, occurred in Hsiaolin village**
- **The main disaster is caused by large-scale landslide.**



## Mitigation Strategies of Large-scale Landslide Disaster in Taiwan

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**Background**

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**Solution**

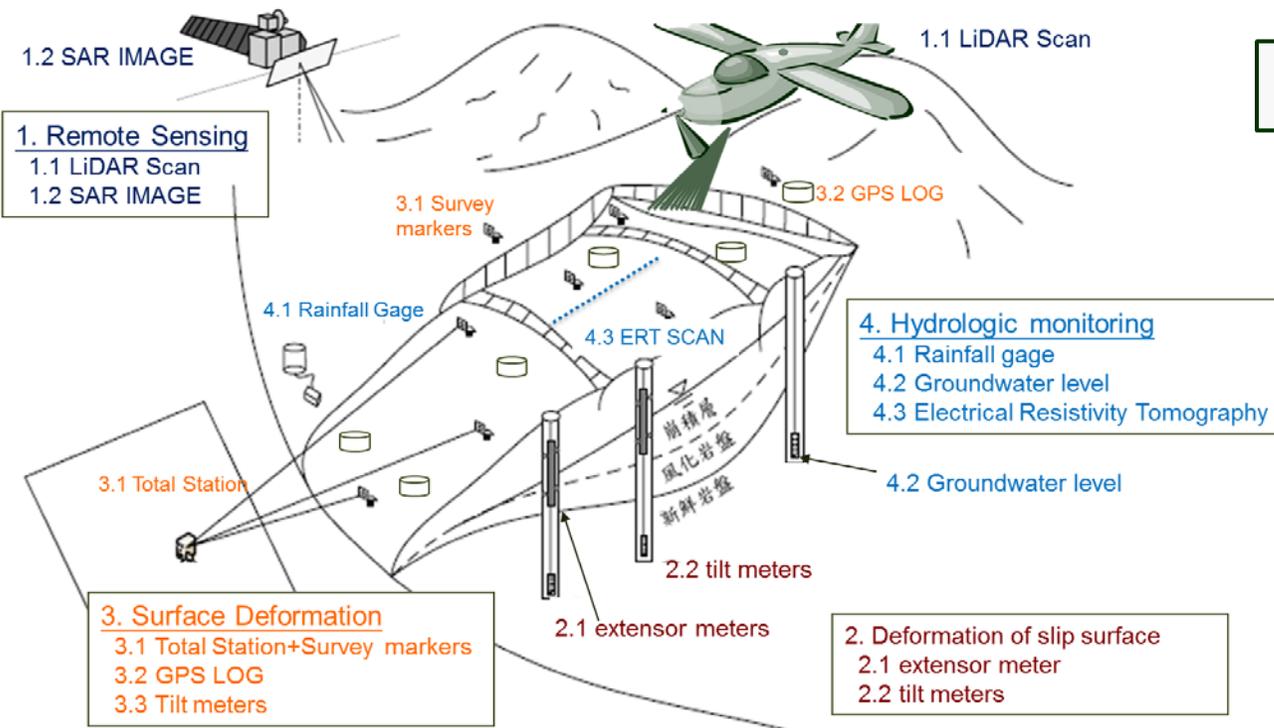
**Comparison**

**Conclusion**

## **IDEA FOR MITIGATION STRATEGIES OF L.S.L. (6Ws1H)**

- **Where** → **Potential Area**
- **What** → **Risk analysis** (Influence Area)
- **Why** → **On-site monitoring**
- **When** → **Forecast and Warning System**
- **Who** → **Evacuation Plan**
- **How** → **Training, Drills, Education**

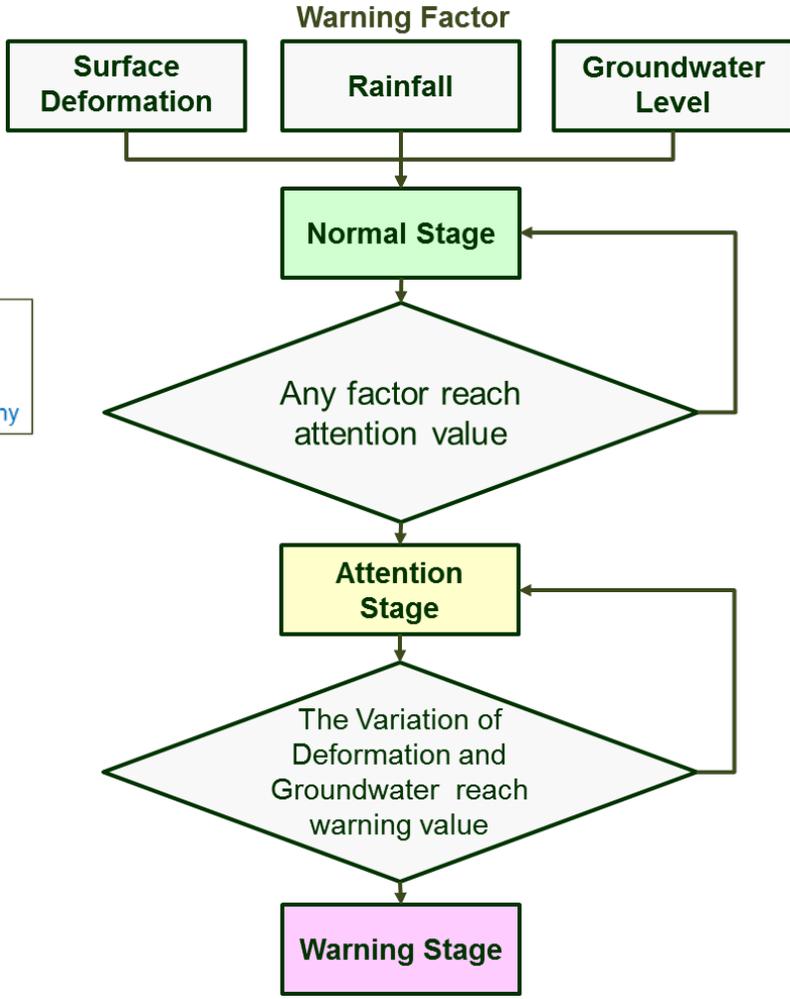
### On-site monitoring



All the monitor result to setup the warning system, such as rainfall, groundwater level, surface deformation.

How to combine all the result together?

### Forecast and Warning System



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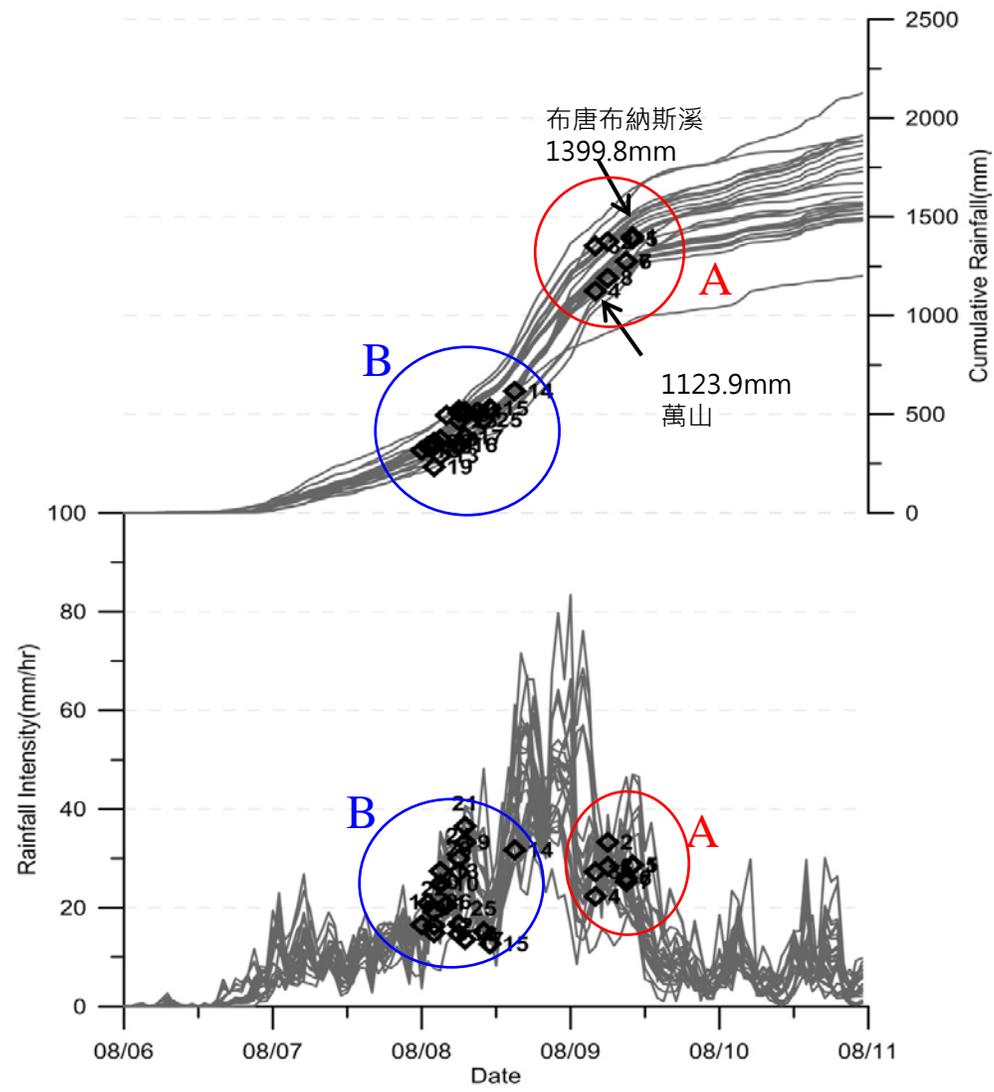
Modelized

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- To Review Landslide cases in typhoon Morakot, 2009
- Group **A** is classified as **new born landslide**, and group **B** include **enlarge case** and **combine case**
- **New Born Case**
  - ❑ Cumulative rainfall >1,000 mm
  - ❑ After rainfall peak
  - ❑ near the turning point of accumulative rainfall
- **Enlarge & Combine case**
  - ❑ Cumulative rainfall is about 300- 500 mm
  - ❑ Before rainfall peak



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# Case Study of Large-scale Landslide in Taiwan

## Definition of Large-scale Landslide

**Area > 10 ha**

**Depth > 10 m**

**Volume >  $10^4$  m<sup>3</sup>**

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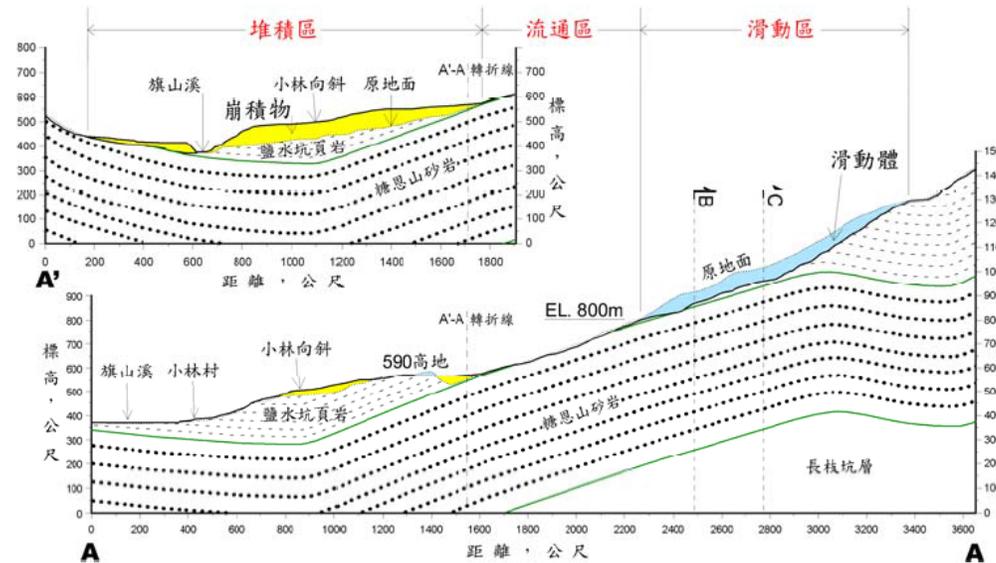
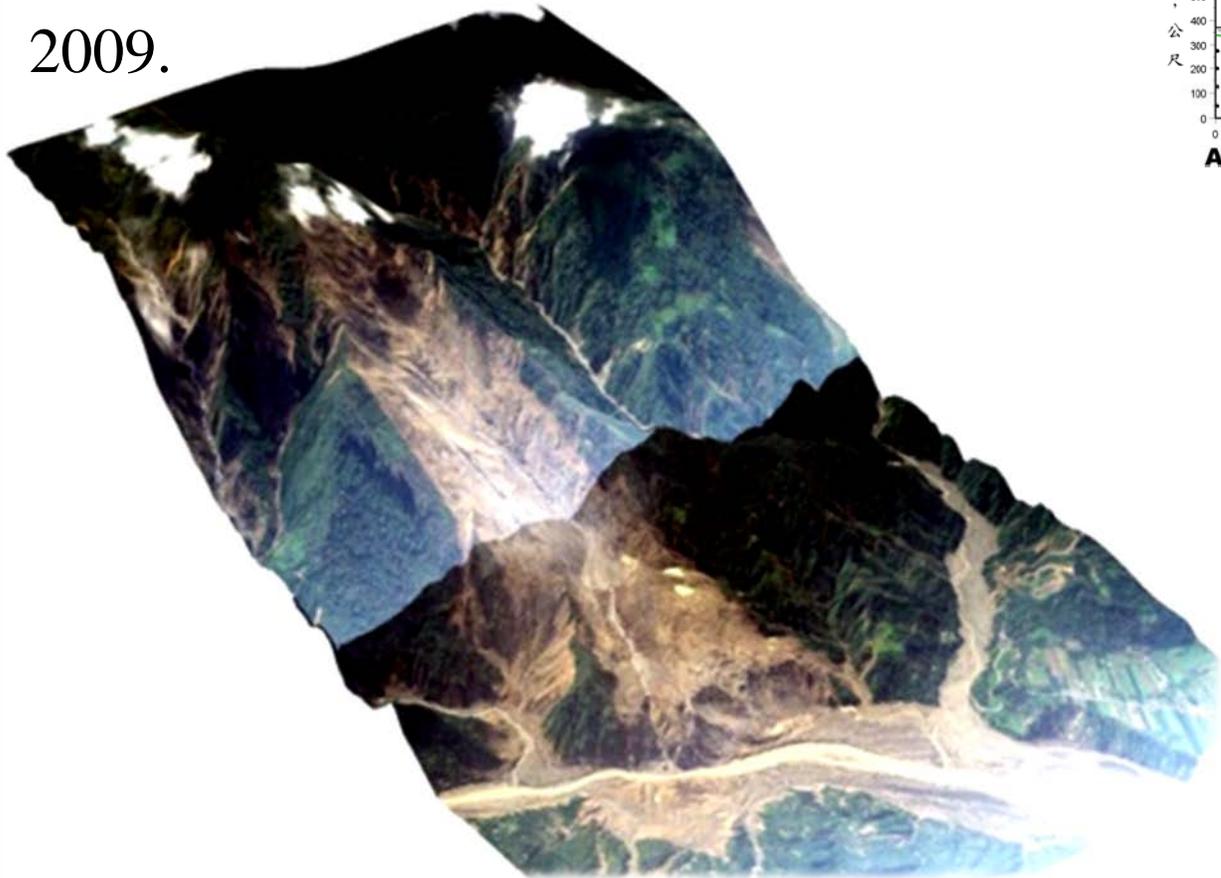
**Solution**

**Comparison**

**Conclusion**

- Fast-moving landslide was **caused by typhoon.**

- Hsiaolin Village, Kaohsiung County in 2009.



圖九 小林村地質剖面圖A-A與A'-A'

- Time : August/06-10 (Typhoon Morakot)
- Area of landslide : 202 ha
- Depth of landslide : 80 m
- **Slope : 26°**
- Movement Distance : 2000 m
- Accumulative Rainfall : 2583 mm
- Loose soil mass
- **With the imprint of groundwater**

- Fast-moving landslide was **caused by typhoon**
- Taimali River, Taitung County in 2009



- **Time : August/06-10 (Typhoon Morakot)**
- **Area of landslide : 290 ha**
- **Depth of landslide : 200 m**
- **Slope : 22°**
- **Movement Distance : 1000 m**
- **Accumulative Rainfall : 1400 mm**
- **Loose soil mass**
- **With the imprint of groundwater**

**Background**

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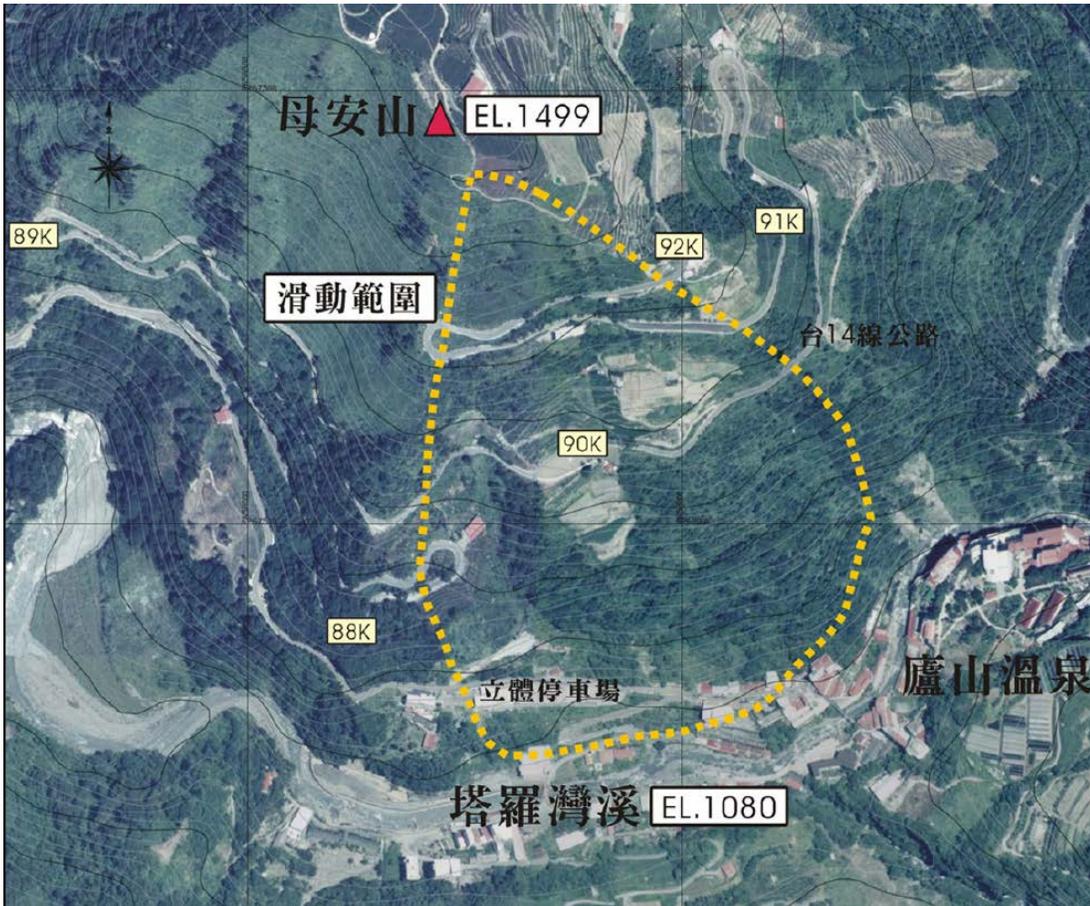
**Modelized**

**Solution**

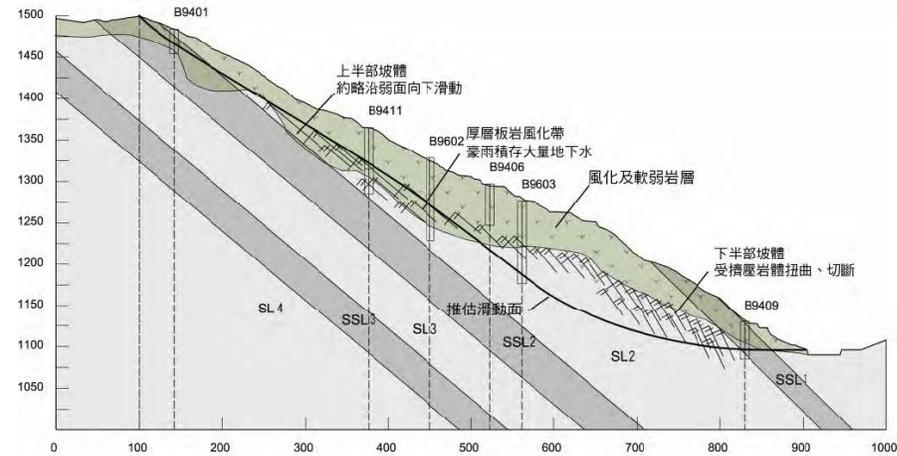
**Comparison**

**Conclusion**

- Slow-moving landslide was **caused by rainfall**
- Lushan landslide, Nantou since 1994



廬山地滑監測及後續治理規劃，黎明工程，2003



- Area of landslide : 30 ha
- Depth of landslide : ~100 m
- **Slope : 26°**
- Mass movement
- **With the imprint of groundwater**

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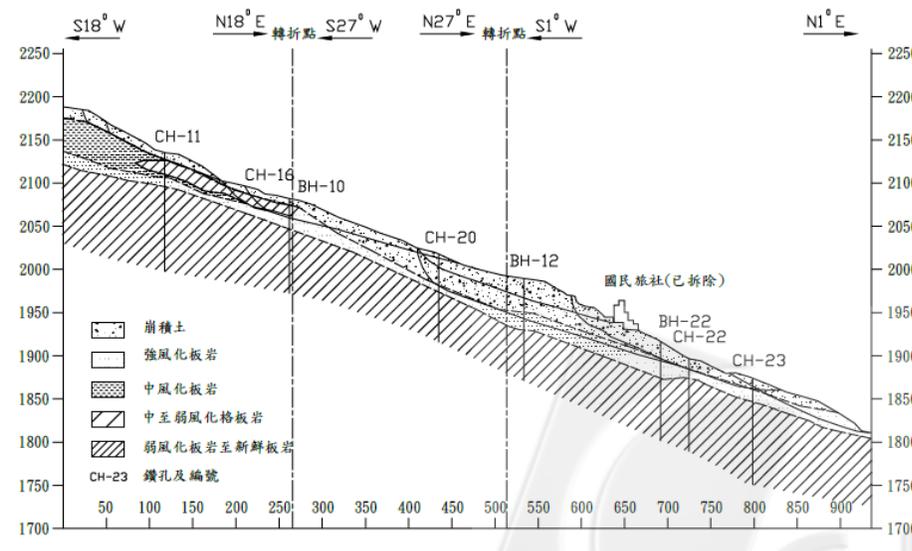
Modelized

Solution

Comparison

Conclusion

- Slow-moving landslide was **caused by rainfall**
- Li-San landslide, Taichung, since 1990



- Area of landslide : 230 ha
- Depth of landslide : ~50 m
- Slope : 20°
- Movement Distance : 800 m
- With the imprint of groundwater

Background

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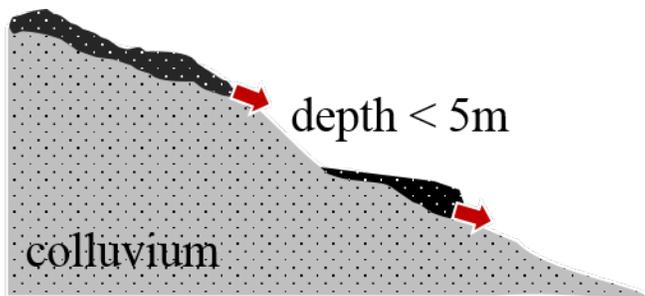
Solution

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# Similarity of F.M.L and S.M.L.

Shallow Landslide  
(fast-moving landslide)



The Experimental Forest, NTU(2012 yr.)

Occurrence: < 2 days(Typhoon Saola)  
without the groundwater layer

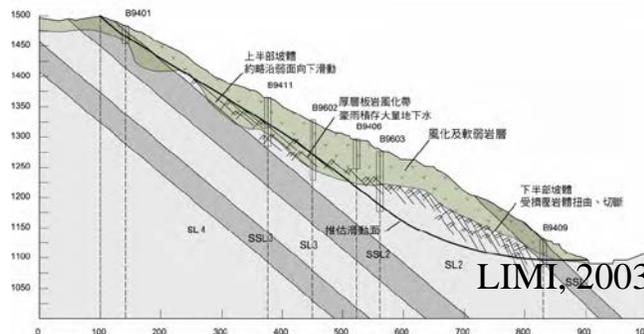
$$\text{Slope} = 22^\circ \sim 29^\circ$$

$$\phi \approx 25^\circ$$

$$i_p = 99.5 \text{ mm/hr}$$

$$P = 1132 \text{ mm}$$

Large-scale Landslide  
(slow-moving landslide)



Lushan landslide

Occurrence: For decades  
with the imprint of groundwater

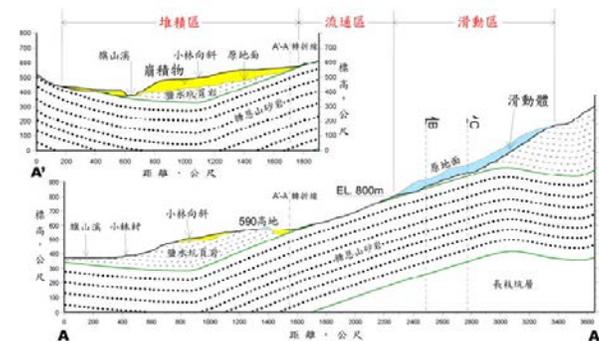
$$\text{Slope} = 24^\circ \sim 26^\circ$$

$$\phi_r \approx 22^\circ \sim 27^\circ$$

$$i_p = 30 \sim 35 \text{ mm/hr}(2005\sim6 \text{ yr.})$$

$$P = 820 \text{ mm} (2006 \text{ yr})$$

Large-Scale Landslide  
(fast-moving landslide)



Occurrence: 4 days  
(Typhoon Morakot)  
with the imprint of groundwater

$$\text{Slope} = 26^\circ$$

$$\phi_r \approx 21^\circ \sim 25^\circ$$

$$i_p = 103 \text{ mm/hr}$$

$$P = 2583 \text{ mm}$$

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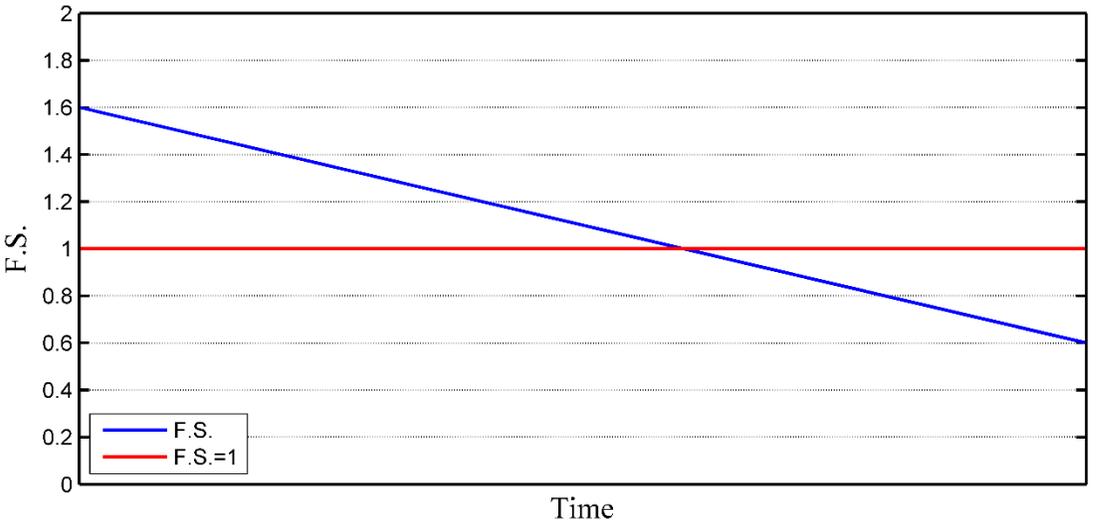
Solution

Comparison

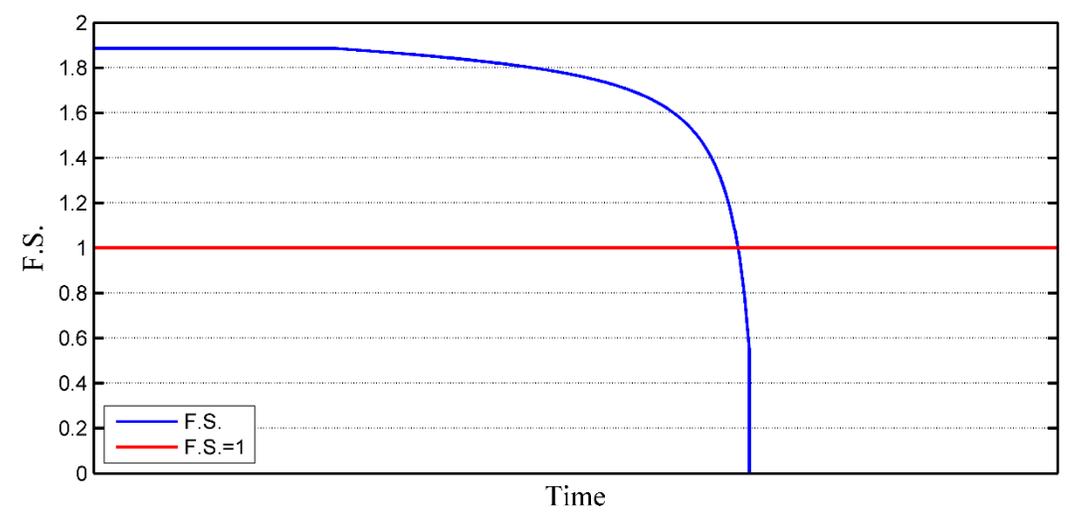
Conclusion

# Difference of F.M.L. and S.M.L.

S.M.L



F.M.L



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Safety Factor  $FS = \frac{R}{D}$

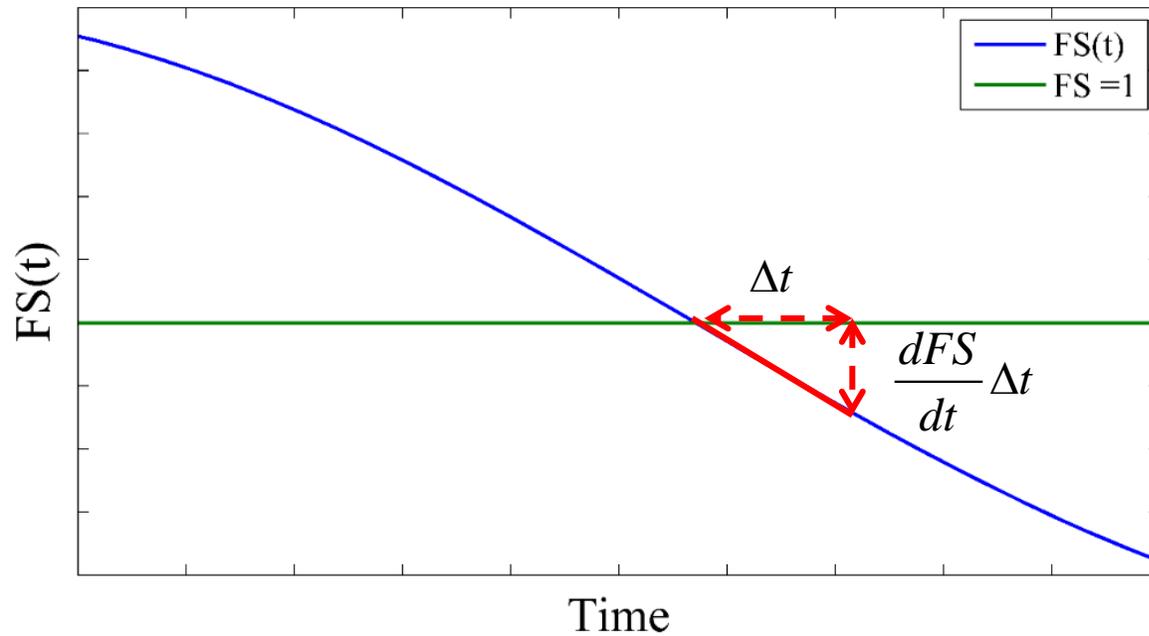
$D - R = (1 - FS)D = \text{Net force acting on sliding block} = Ma$

$\Rightarrow \frac{D}{M}(1 - FS) = a, \text{ where } 1 - FS = \frac{dFS}{dt} \Delta t$

$\Rightarrow a = \frac{D}{M} \frac{dFS}{dt} \Delta t$

$\Rightarrow a \propto \frac{dFS}{dt}$

D: Driving Force  
R: Resistance Force  
M: Mass of sliding block



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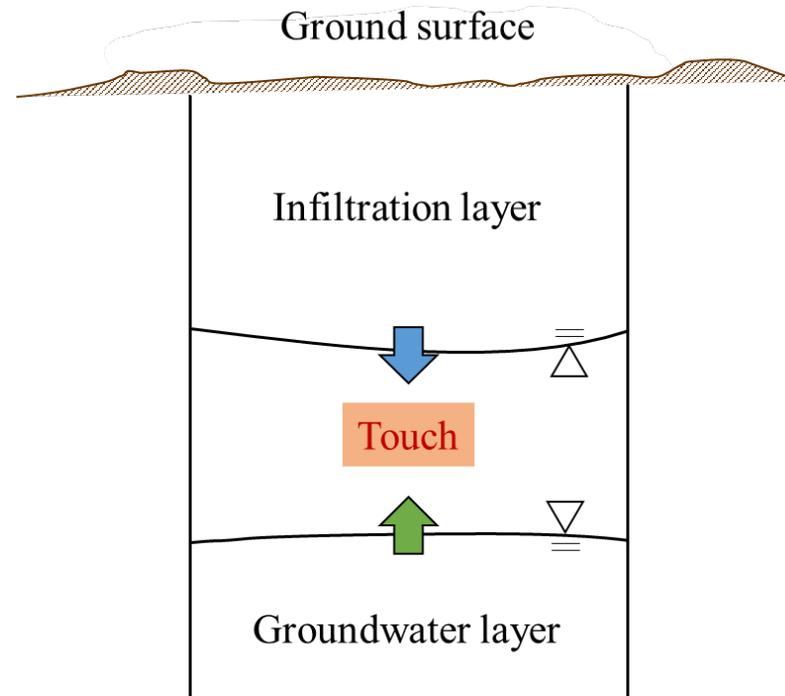
**Modelized**

**Solution**

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- The flow velocity of infiltration and groundwater is usually small, therefore, what's reason causes the rapid decreasing in F.S.?
- We are interesting in **what kinds of interaction between infiltration and groundwater** will trigger sudden failure or creeping landslide.



**Background**

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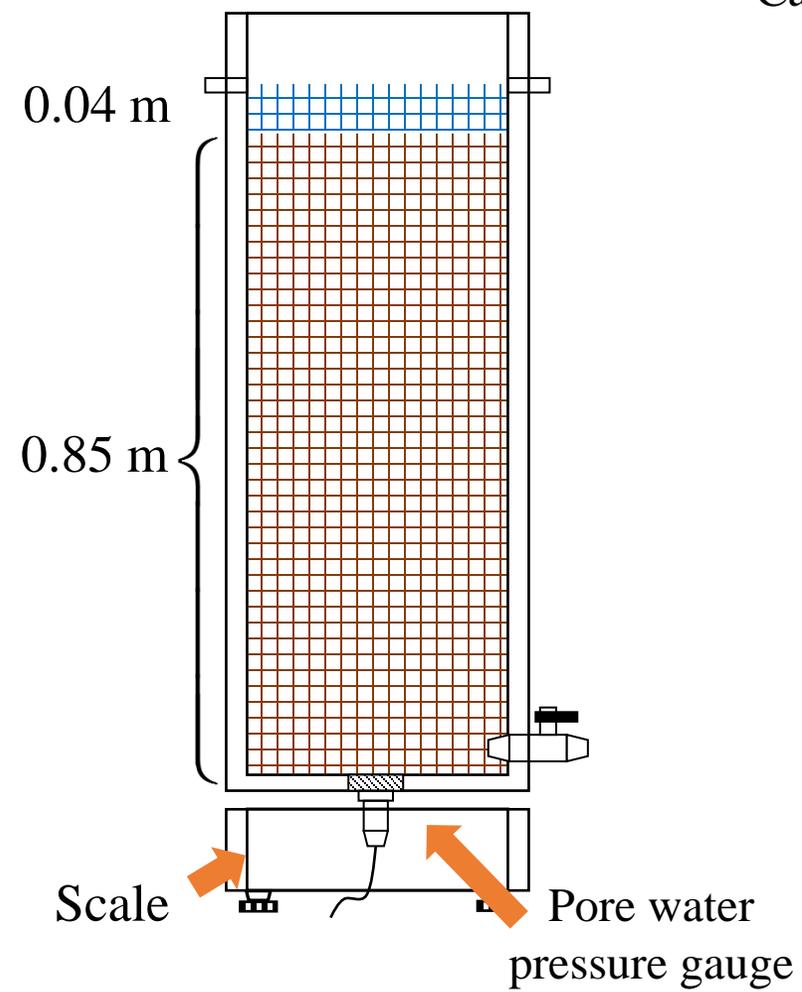
**Experiment**

**Modelized**

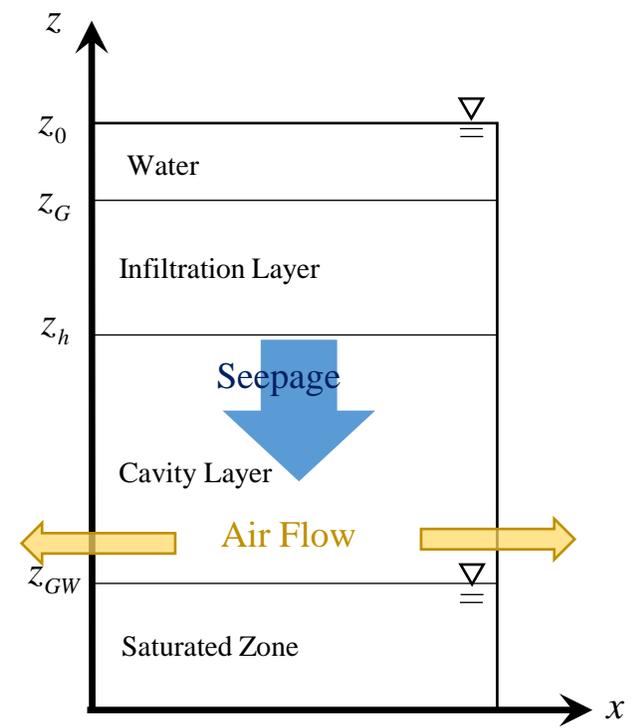
**Solution**

**Comparison**

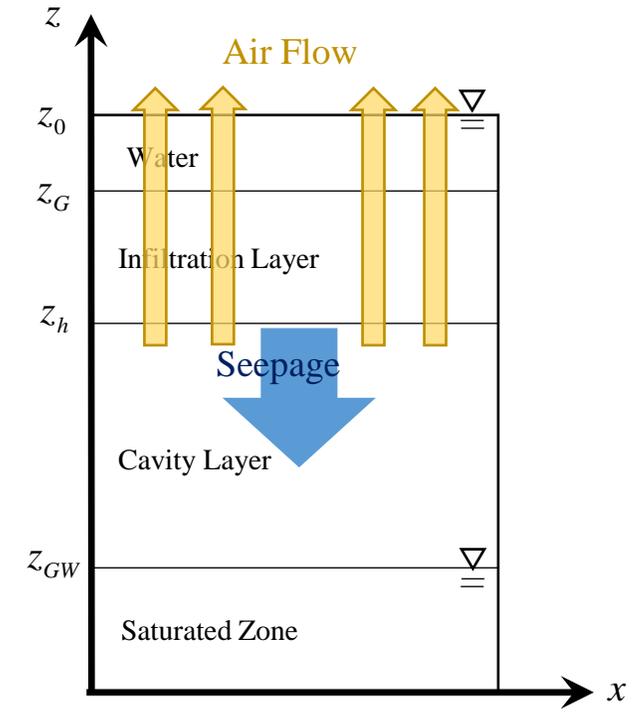
**Conclusion**



Case I : Air escape from bottom



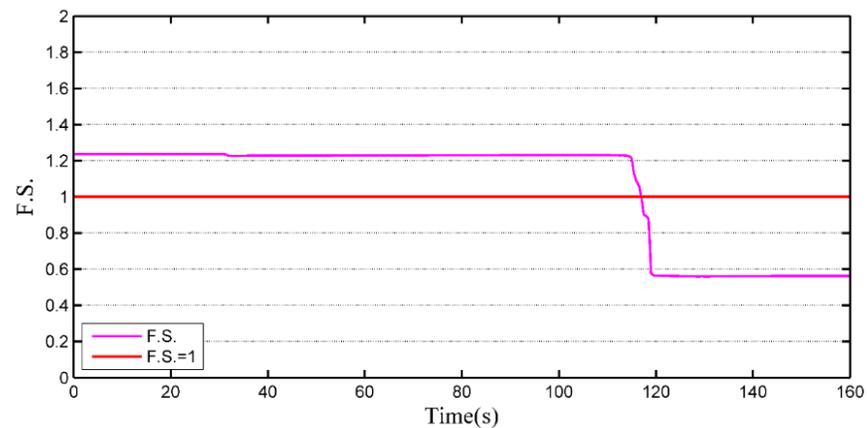
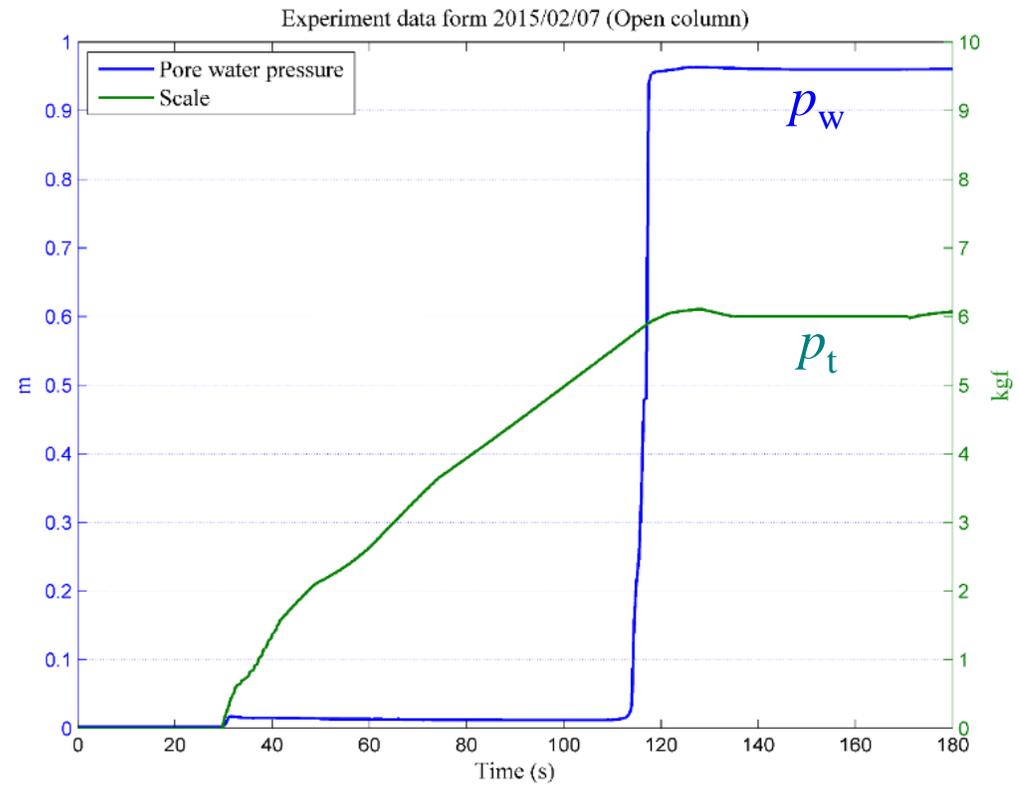
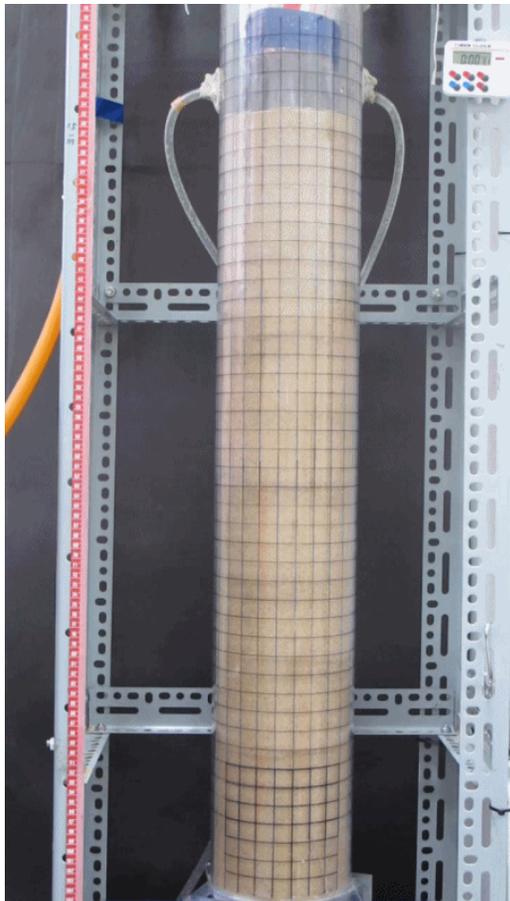
Case III : Air escape from top



Experiment measure data  $\left\{ \begin{array}{l} W_{sc} : \text{weight} \\ p_{wb} : \text{porewater pressure} \end{array} \right.$

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**Case I: Air escape from bottom**



$$S.F. = \frac{c + \{(1-n)\gamma_s L_s + Sn\gamma_w h(t) - p_w(t)\} \tan \varphi \cos \theta}{\{(1-n)\gamma_s L_s + Sn\gamma_w h(t)\} \sin \theta}$$

where,  $\varphi = 25^\circ; \theta = 30^\circ; c = 0.$

**Background**

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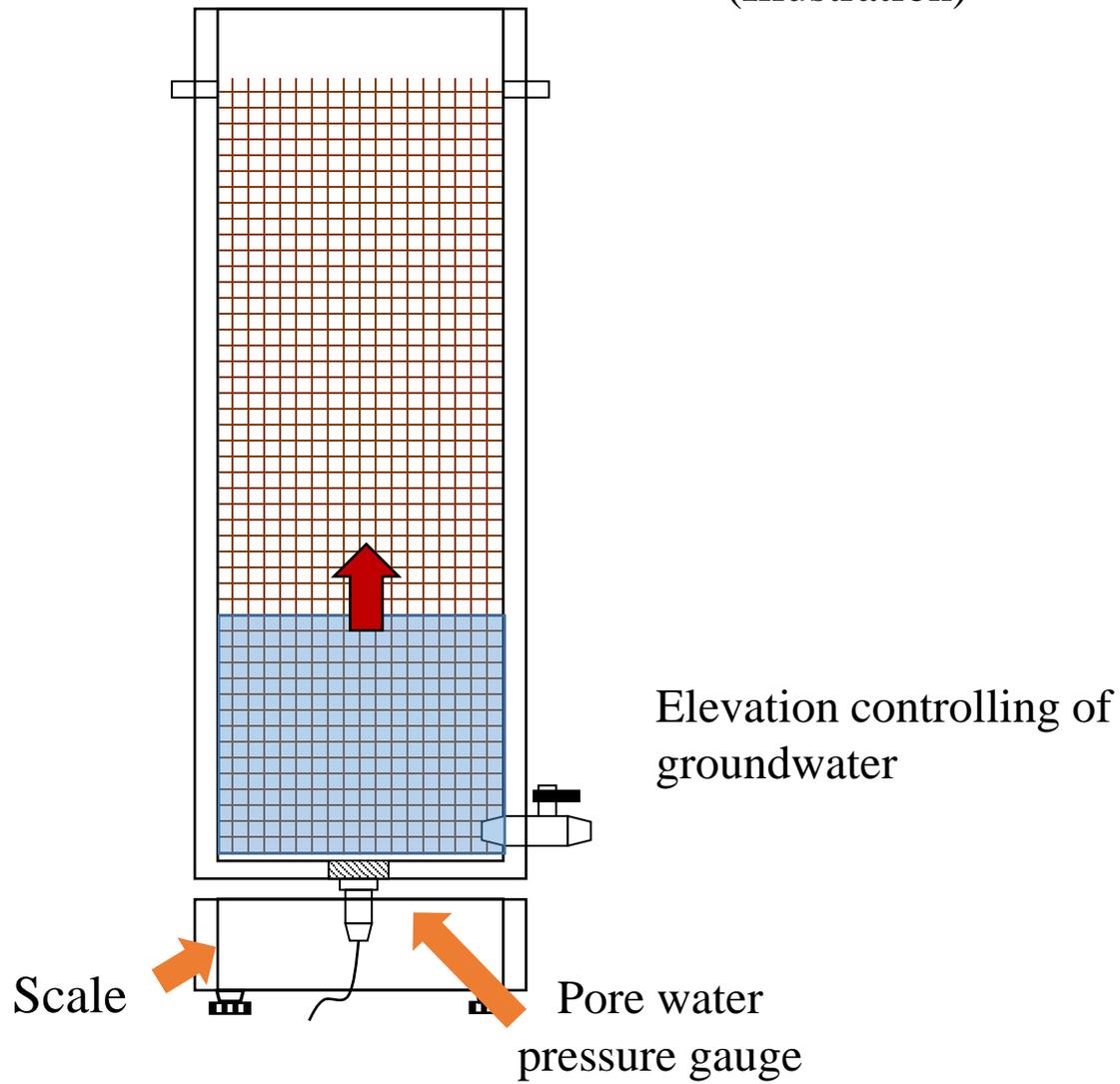
**Solution**

**Comparison**

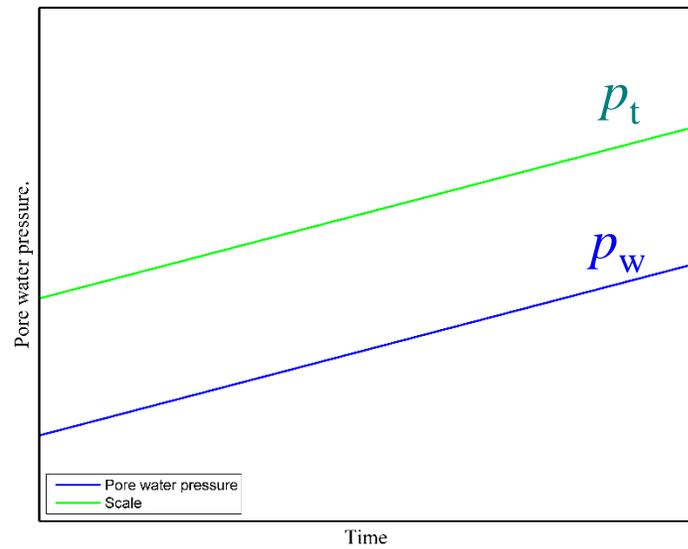
**Conclusion**

**Case II: Groundwater upwelling**

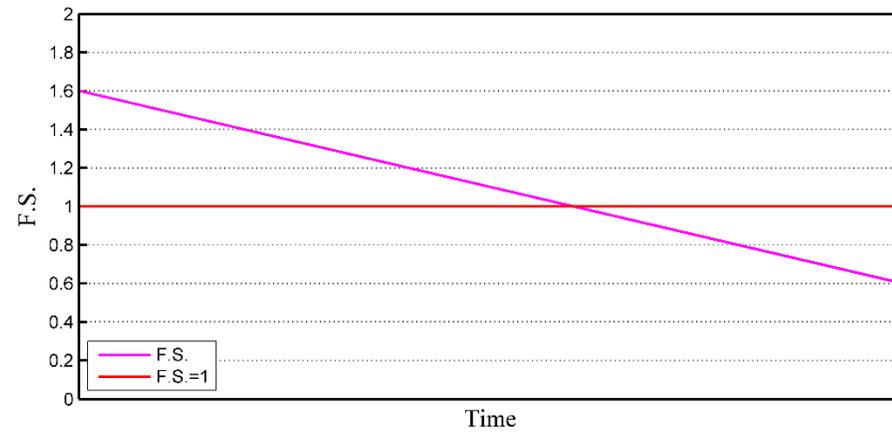
(Illustration)



Anticipation of change in pore water pressure



Anticipation of change in S.F.



**Background**

**Case Study**

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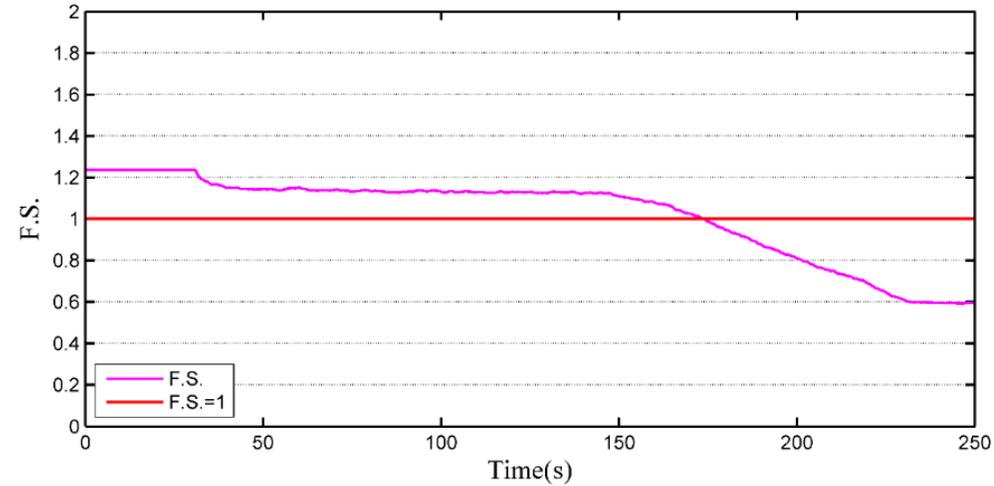
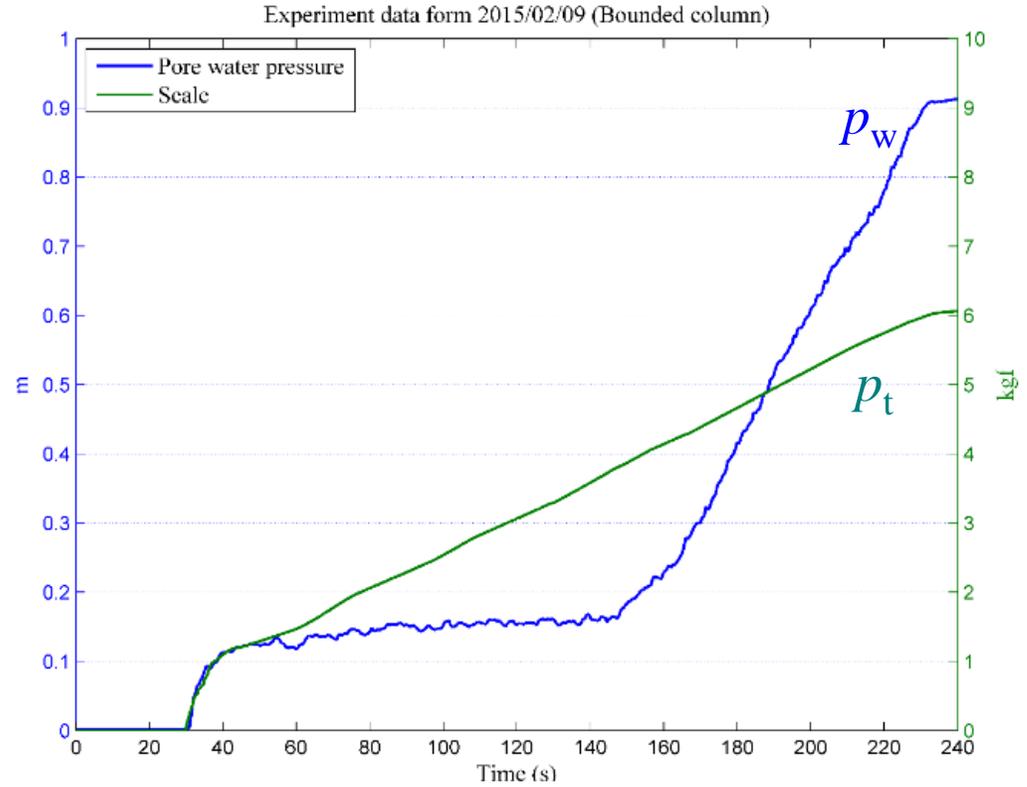
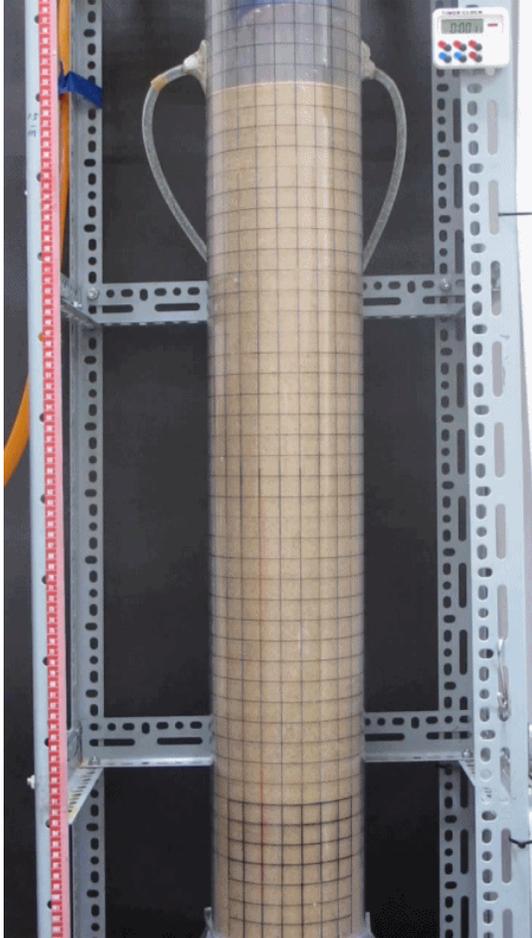
**Modelized**

**Solution**

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**Case III: Air escape from top**



$$S.F. = \frac{c + \{(1-n)\gamma_s L_s + Sn\gamma_w h(t) - p_w(t)\} \tan \phi \cos \theta}{\{(1-n)\gamma_s L_s + Sn\gamma_w h(t)\} \sin \theta}$$

where,  $\phi = 25^\circ$ ;  $\theta = 30^\circ$ ;  $c = 0$ .

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- Develop the theory from physical phenomenon

Equation of mass conservation

$$\frac{\partial nS \rho_w}{\partial t} + \nabla \cdot nS \rho_w \mathbf{v}_w = 0$$

$$\frac{\partial n(1-S) \rho_a}{\partial t} + \nabla \cdot n(1-S) \rho_a \mathbf{v}_a = 0$$

Equation of momentum conservation

$$nS \rho_w \frac{D\mathbf{v}_w}{Dt} = nS \rho_w \mathbf{g} + \nabla \cdot (-p_w \mathbf{I}) + \mathbf{f}_w$$

$$n(1-S) \rho_a \frac{D\mathbf{v}_a}{Dt} = n(1-S) \rho_a \mathbf{g} + \nabla \cdot (-p_a \mathbf{I}) + \mathbf{f}_a$$

$$(1-n) \rho_s \frac{D\mathbf{v}_s}{Dt} = (1-n) \rho_s \mathbf{g} + \nabla \cdot \boldsymbol{\sigma}_s^* + \mathbf{f}_s$$

Equation of state

$$p_a V = \frac{m}{M} \bar{R} T$$

Interaction among the constituents

$$\mathbf{f}_w + \mathbf{f}_s + \mathbf{f}_a = 0$$

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• Convert the 3-D governing equation into 1-D.

• Interaction among the constituents in 1-D

-the momentum supplies(interactions) can be divided into two parts

$f^{buoy}$ : caused by the **normal pressure** in the constituents surface

$f^{rel}$ : the drag force due to the **relative velocity** tangential to the constituents surface

$$f_w$$

$$f_w^{buoy} = (1-n) \frac{nS}{n} \frac{\partial p_w}{\partial z} + \frac{n(1-S)}{n} \frac{\partial p_w}{\partial z}$$

$$f_w^{rel} = -nS \frac{\partial p_{iw}}{\partial z}$$

$$f_s$$

$$f_s^{buoy} = - \left\{ (1-n) \frac{nS}{n} \frac{\partial p_w}{\partial z} + (1-n) \frac{n(1-S)}{n} \frac{\partial p_a}{\partial z} \right\}$$

$$f_s^{rel} = nS \frac{\partial p_{iw}}{\partial z} + n(1-S) \frac{\partial p_{ia}}{\partial z}$$

$$f_a$$

$$f_a^{buoy} = (1-n) \frac{n(1-S)}{n} \frac{\partial p_a}{\partial z} - \frac{n(1-S)}{n} \frac{\partial p_w}{\partial z}$$

$$f_a^{rel} = -n(1-S) \frac{\partial p_{ia}}{\partial z}$$

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# • Governing equation in 1-D form

Equation of mass conservation

$$\frac{\partial}{\partial t}(nS\rho_w) + \frac{\partial}{\partial z}(nS\rho_w u_w) = 0$$

$$\frac{\partial}{\partial t}\{n(1-S)\rho_a\} + \frac{\partial}{\partial z}\{n(1-S)\rho_a u_a\} = 0$$

Equation of momentum conservation

$$nS\rho_w \frac{Du_w}{Dt} = -nS\rho_w g - nS \left( \frac{\partial p_w}{\partial z} + \frac{\partial p_{iw}}{\partial z} \right)$$

$$n(1-S)\rho_a \frac{Du_a}{Dt} = -n(1-S)\rho_a g + (nS - S - n) \frac{\partial p_a}{\partial z} - (1-S) \frac{\partial p_w}{\partial z} - n(1-S) \frac{\partial p_{ia}}{\partial z}$$

$$(1-n)\rho_s \frac{Du_s}{Dt} = -(1-n)\rho_s g - \frac{\partial p_s}{\partial z} + S \left\{ n \frac{\partial p_{iw}}{\partial z} - (1-n) \frac{\partial p_w}{\partial z} \right\} + (1-S) \left\{ n \frac{\partial p_{ia}}{\partial z} - (1-n) \frac{\partial p_a}{\partial z} \right\}$$

Equation of state

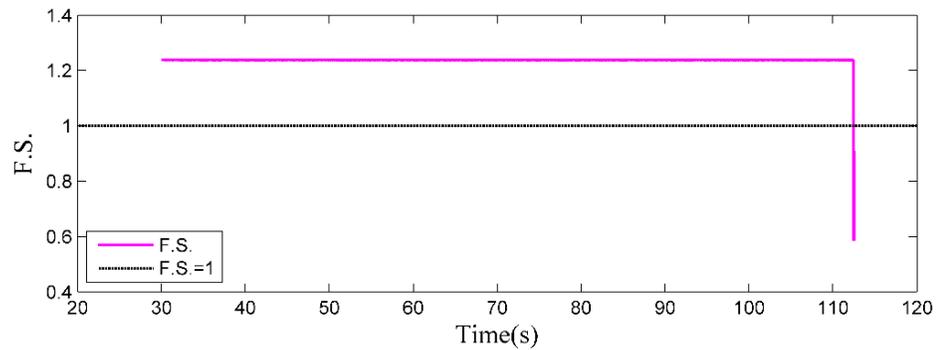
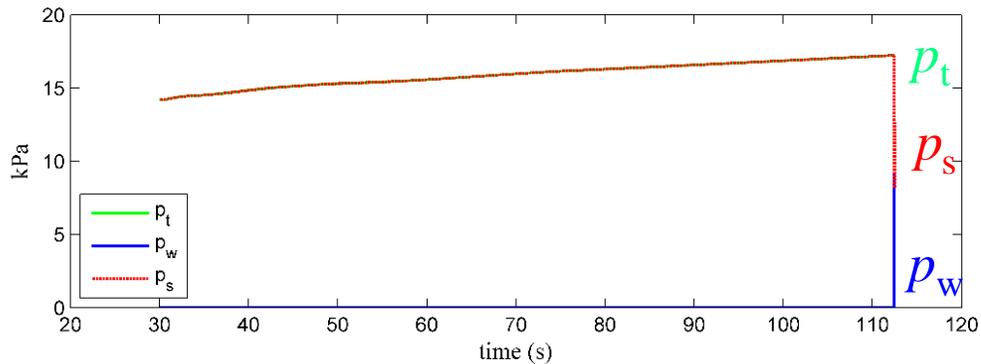
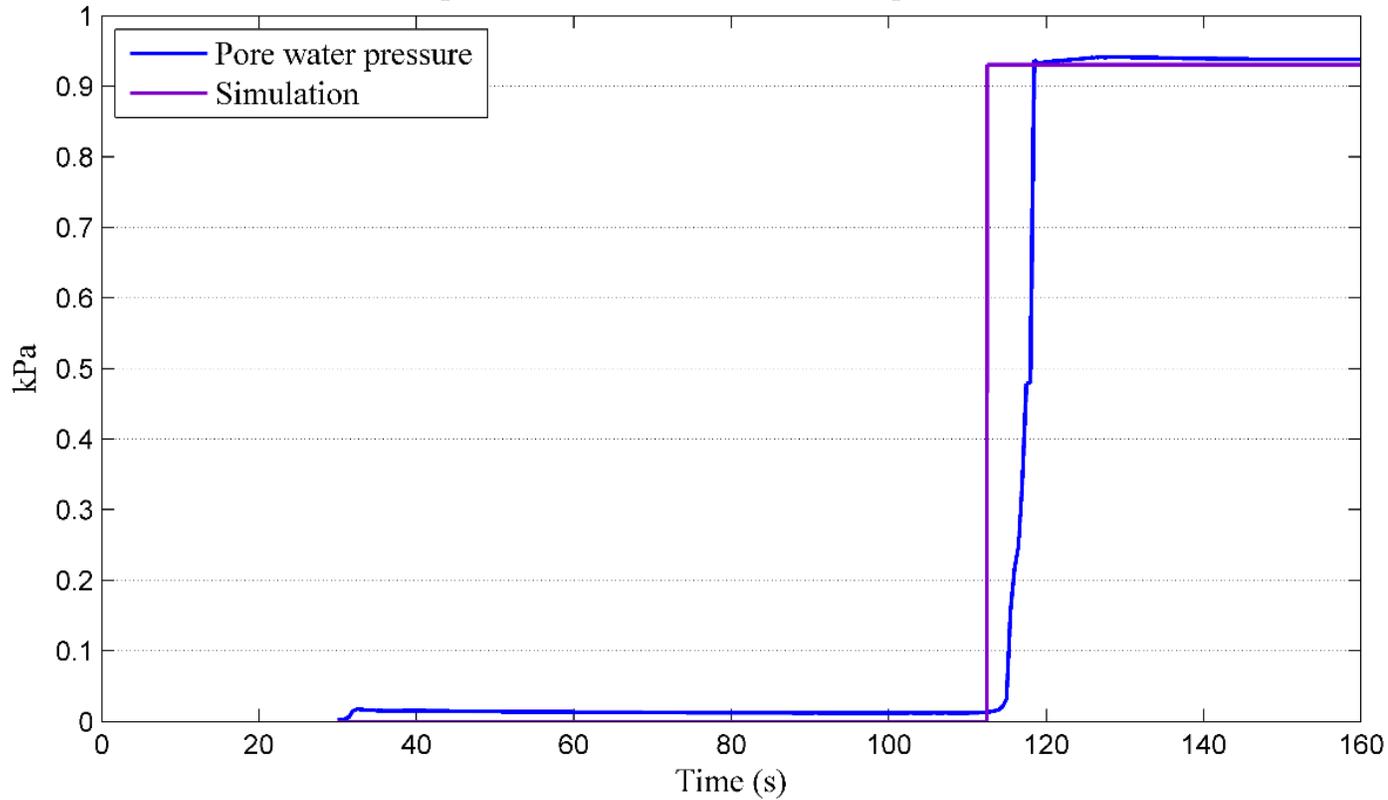
$$p_a z A = \frac{m}{M} \bar{R} T$$

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### Case I: Air escape from bottom

Experiment data form 2015/02/07 (Open column)



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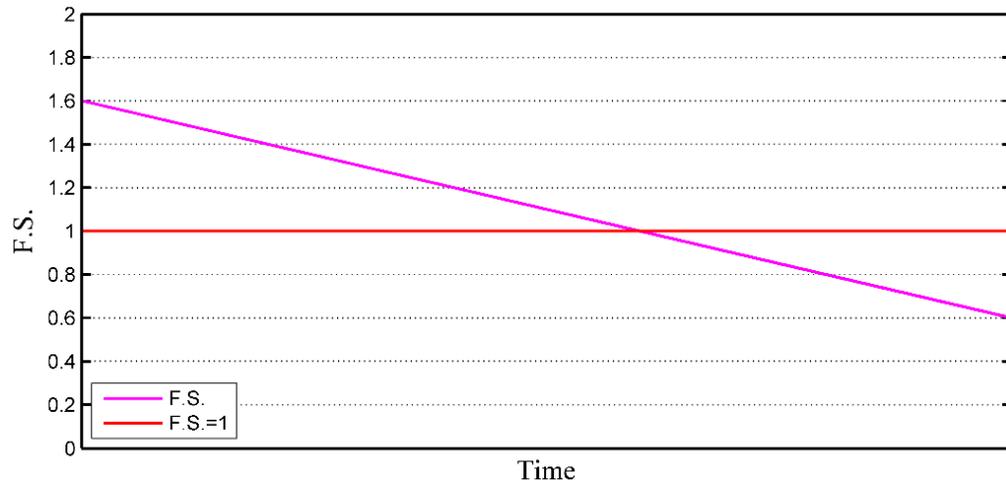
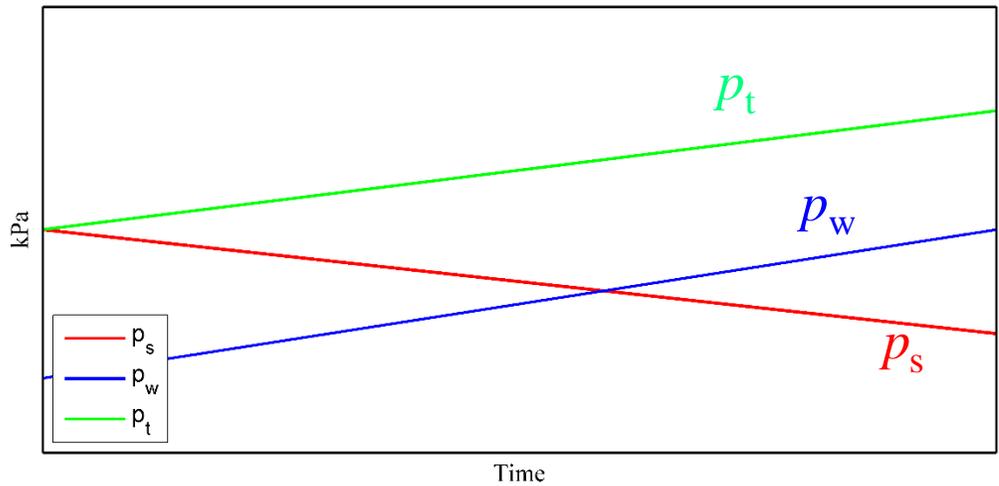
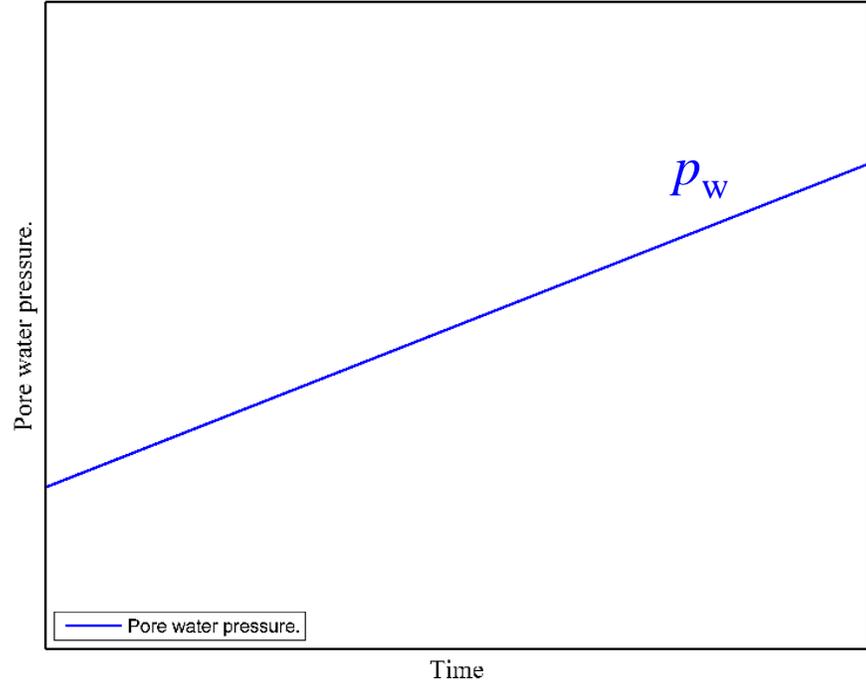
Solution

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**Case II: Groundwater upwelling**



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**Solution**

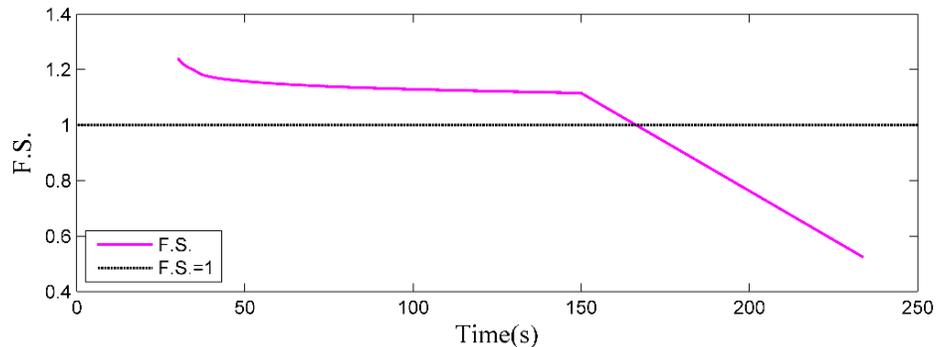
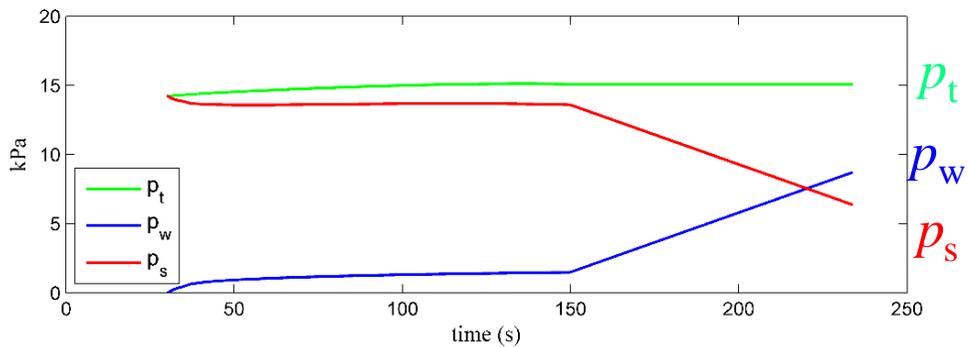
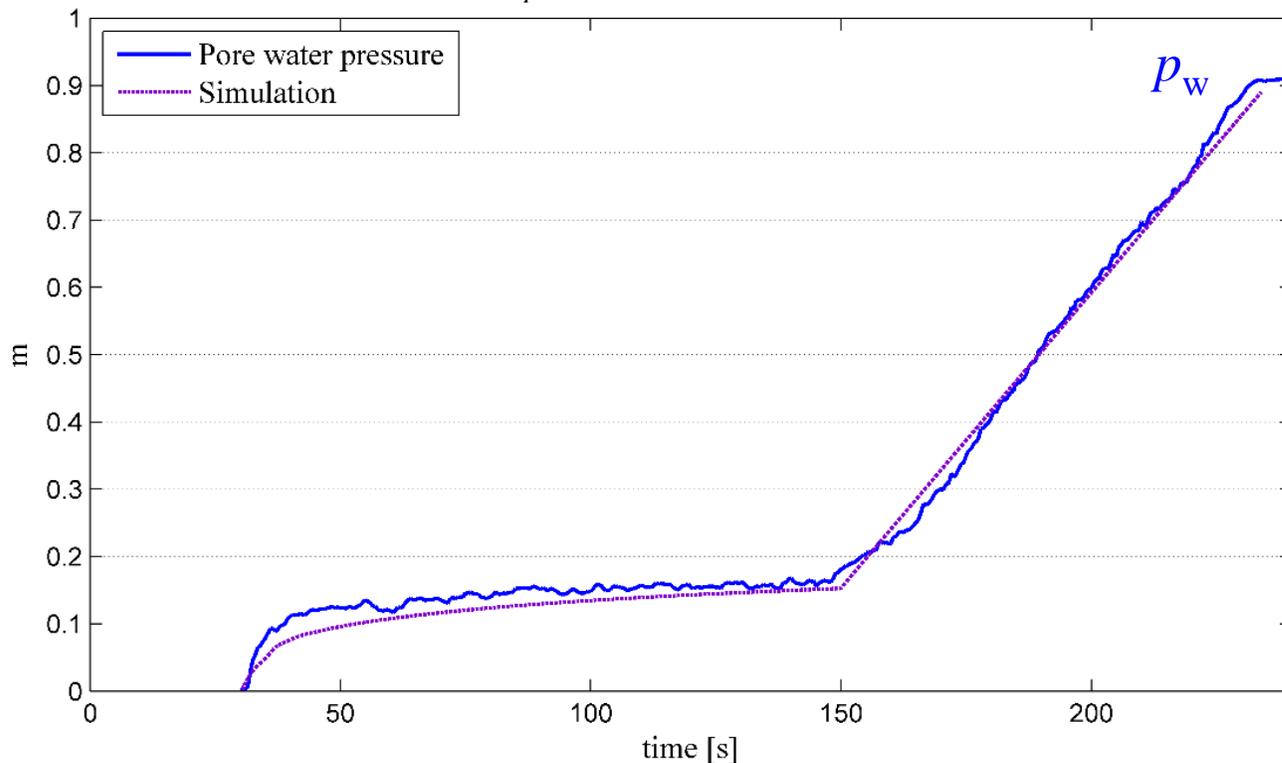
**Comparison**

**Conclusion**



### Case III: Air escape from top

Experiment data form 20150209



**Background**

**Case Study**

**Idea**

**Experiment**

**Modelized**

**Solution**

**Comparison**

**Conclusion**



# Conclusion

**Background**

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- This model is a pilot study of LSL warning system in Taiwan . More field data are necessary before it is used.
- The moment of infiltration front touch with groundwater will induce fast-moving LSL.

**Background**

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# *Don't Touch !!*



**Background**

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